

Is eucalyptol the cause of nocturnal events observed in Australia?

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Nocturnal nucleation events were observed from July 2005 to October 2006 in Tumbarumba in New South Wales, Australia. These events were observed on one third of the nights and they were often much more intense than normal daytime events. One of the main features found in this environment was the abundance of eucalyptol among the volatile organic compounds (VOC) emitted by the local Eucalypt forest. In contrast, in most other forest environments, such as Hyttiälä in southern Finland, eucalyptol is a minor component of VOC emissions. Therefore, the objective of this study was to evaluate the role of eucalyptol in the nocturnal nucleation events. For this purpose, a series of experiments and quantum mechanical calculations were performed. Both approaches showed that the role of eucalyptol in the nocturnal events can be ruled out.

Introduction

Aerosols are ubiquitous in the Earth's lower atmosphere. They affect human health, visibility, atmospheric chemistry and climate. Aerosols influence climate directly by scattering and absorbing radiation, and indirectly by acting as cloud condensation nuclei and affecting cloud properties. Gas-to-particle nucleation is an important source of new aerosol particles in the Earth's atmosphere (Kulmala *et al.* 2004a). This phenomenon has been the subject of intense studies in the last years, but the formation mechanism and the participating substances have not

been resolved yet. A strong correlation has been observed between new-particle formation and sulfuric acid concentrations (Weber *et al.* 1996, Weber *et al.* 1997, Kulmala *et al.* 2000, Sihto *et al.* 2006, Riipinen *et al.* 2007). However, in many cases the observed rates of particle formation are much higher than the rates expected for binary homogeneous sulfuric acid-water nucleation. Different additional processes have been proposed to explain this difference, such as homogeneous ternary water-sulfuric acid-ammonia nucleation (Ball *et al.* 1999, Korhonen *et al.* 1999, Anttila *et al.* 2005) participation of organic molecules (Yu *et al.* 1999, Hoffmann *et al.* 1998)

and ion induced nucleation (IIN) (Lovejoy *et al.* 2004, Yu and Turco 2000 and 2008, Boy *et al.* 2008). The contribution of IIN to the total number of particle formed is still under discussion. Examples are the studies by Yu and Turco (2008), Kulmala *et al.* (2007) and Boy *et al.* (2008) carried out at the same measuring site, the first one concluded that the contribution of IIN to new formed particle is close to 100% while the second and third studies concluded that the contribution of INN is less than 10%. This shows how uncertain our ability is to quantify the contribution of one process to the total amount of newly formed particles. On the other hand, the recently suggested activation theory (Kulmala *et al.* 2006) is able to explain atmospheric nucleation relatively well.

Organic volatile compounds (VOCs) have been suggested to contribute to the growth of nucleated clusters (O'Dowd *et al.* 2002, Kulmala *et al.* 2004b), although a recent work by Bonn *et al.* (2008) proposes that the VOCs can also be implied in the formation of the clusters from the very first steps.

Nucleation events are in most cases observed during daytime but in some locations they can occur also during the night. In Tumbarumba in New South Wales, Australia, these nocturnal events are very intense and frequently observed (32% of the nights in 2006; Suni *et al.* 2008). If we compare Tumbarumba with other sites where the nocturnal events have been observed but where they are much less intense and less frequent, like Hyytiälä in southern Finland (Junninen *et al.* 2008), one of the main differences turns out to be the composition of the forest. The forest in Tumbarumba is composed mainly of eucalypt trees and that in Hyytiälä of Scots pine. As a result, also the emitted VOCs are different. In Tumbarumba the most abundant monoterpenes was eucalyptol (1,8-cineol), whereas in Hyytiälä carene, α - and β -pinene were emitted the most. Another clear difference between these two environments is the higher concentration of radon found in Tumbarumba.

The objective of this study was to evaluate the possible role of eucalyptol in nocturnal new particle formation. We designed a series of chamber experiments with the objective to reproduce the nocturnal events observed in Aus-

tralia. Additionally, we used quantum chemical calculations to study the molecular mechanism behind the nocturnal nucleation events.

Material and methods

Experimental setup

We conducted the experiments in a 14-m³ isolated room that was ventilated with an air flow of 40 liters per minute. The flow could optionally be piped through a small tank which contained a radon source to introduce up to 2000 Bq m⁻³ of radon in the room. The radon (²²²Rn) concentration was continuously measured using a radon monitor Alphaguard P 30 (Genitron Instruments). Temperature and relative humidity conditions inside the room were monitored using Delta OHM DO 9847 logger with capacitive Mk-33 and Pt100 sensors.

We placed an Air Ion Spectrometer (AIS) inside the chamber to obtain the number size distribution of negative and positive air ions. The instrument is designed to make simultaneous ion distribution measurements of both polarities over a mobility range of 3.2–0.0013 cm² V⁻¹ s⁻¹. The mobility distribution is divided to 27 logarithmically uniform bins which correspond to a Millikan diameter range of 0.8–40 nm. Mirme *et al.* (2007) give the details of this instrument and data inversion.

O₃ and SO₂ concentrations inside the room were measured with a Dasibi Model 1008-RS UV-absorption ozone photometer and a pulsed fluorescence SO₂ analyzer model 43C (Thermo Environmental Instruments, Inc.). A small piece of seal on the side of the chamber window was removed to allow closing the window and yet leading the inlet tubes (PTFE-Teflon) inside the chamber through the hole in the seal.

The concentration of volatile organic compounds (eucalyptol) was determined using Tenax-TA/carbobackB adsorbent tubes. We pumped air from the room through the Tenax tubes for 10 minutes with a constant flow of 100 ml per minute. The samples were analyzed afterwards in a laboratory using an automatic thermodesorption device connected to a gas chromatograph (HP-5890) and a mass spectrometer (HP-5972).

Theoretical method

Our calculations were performed using a systematic multi-step method recently developed by our group. This method is described elsewhere (Ortega *et al.* 2008), so only the relevant details are given here.

The multi-step method allows the study of clusters containing a large organic molecule and up to four sulfuric acid molecules using a reasonable amount of computing time. In this study, the initial guess geometries were chosen by using chemical intuition. These structures were pre-optimized at a low level of theory using SPARTAN program (<http://wavefun.com>). Then the more stable isomers among all initial guesses (usually between 3 and 4) were optimized with the SIESTA program (Soler *et al.* 2002), using the gradient-corrected BLYP functional (Miehlich *et al.* 1989) and the double- ζ polarized (DZP) functions. Vibrational harmonic frequencies were also calculated using SIESTA and were used to estimate the entropy and thermal contributions to the enthalpy and Gibbs free energy of the clusters.

To obtain a reliable energy, the optimized structures obtained from the SIESTA program were used to perform high-level, single-point energy calculations by using the TURBOMOLE program (Ahlrichs *et al.* 1989) with the Resolution of Identity-Coupled Cluster Single and Doubles method (RI-CC2) (Hättig *et al.* 2000) and the aug-cc-pV($T+d$)Z (Dunning *et al.* 2001) basis set, which is identical to aug-cc-pVTZ for hydrogen, oxygen and nitrogen atoms, and contains one extra set of d -orbitals for the sulfur atoms.

The calculated values for entropy and thermal contributions to the enthalpy and Gibbs free energy (ΔG) of the clusters were calculated assuming a pressure of 1013 hPa for the monomers. In order to obtain a realistic picture of the free energies in atmospheric conditions, we have to take into account the relative concentration of each molecular species in atmosphere. This was done via the law of mass action as showed in Ortega *et al.* (2008).

Results and discussion

Experimental results

We carried out different repetitive chamber experiments all leading to the same results, so only two examples are shown. In the first example we used only eucalyptol in the chamber, whereas in the second one we used eucalyptol and radon to simulate the high concentration of radon found in Tumbarumba.

Eucalyptol experiments

We carried out the first experiment on 7 April 2008. We used 1 ml of 99% pure eucalyptol (Sigma-Aldrich) in a glass container cooled with ice in order to decrease the vapor pressure of eucalyptol and keep the amount of eucalyptol introduced into the chamber as low as possible.

Before introducing the eucalyptol in the chamber, we took two Tenax tube samples to check the background concentration of organics in the chamber. The concentrations of all organic compounds measured were below 0.6 ng l⁻¹.

The eucalyptol was introduced in the chamber at 11:12. After four hours no event was observed (Fig. 1). Three Tenax tube samples were taken during that period (Table 1).

Within the first half hour we got a burst of eucalyptol, and after one hour its concentration dropped to around 10 ng l⁻¹ (4×10^{10} cm⁻³) at which level it appeared to remain. The highest concentration of eucalyptol found in Tumbarumba was 0.29 ng l⁻¹ (1.13×10^9 cm⁻³) during the night, so the concentration in our experiments was around 30 times higher.

The conditions after the eucalyptol was introduced in the chamber were quite constant (Fig. 2), with the relative humidity between 27% and 31%, temperature between 26 and 29 °C, SO₂ concentration around 0.2 ppb (4.93×10^9 cm⁻³) and ozone concentration between 2 and 14 ppb (1.97×10^{11} cm⁻³).

In Tumbarumba, ozone and SO₂ were only measured during an intensive campaign in November 2006 (T. Suni unpubl. data). At that time, the ozone concentration ranged from 10 to

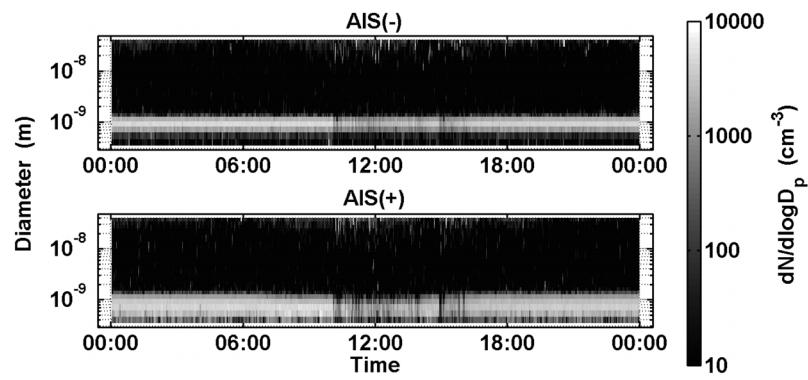


Fig. 1. Surface plot of air ion distribution measured on 7 April 2008 over size range of 0.34–40 nm in the Tammett diameter (Tammet *et al.* 1995)

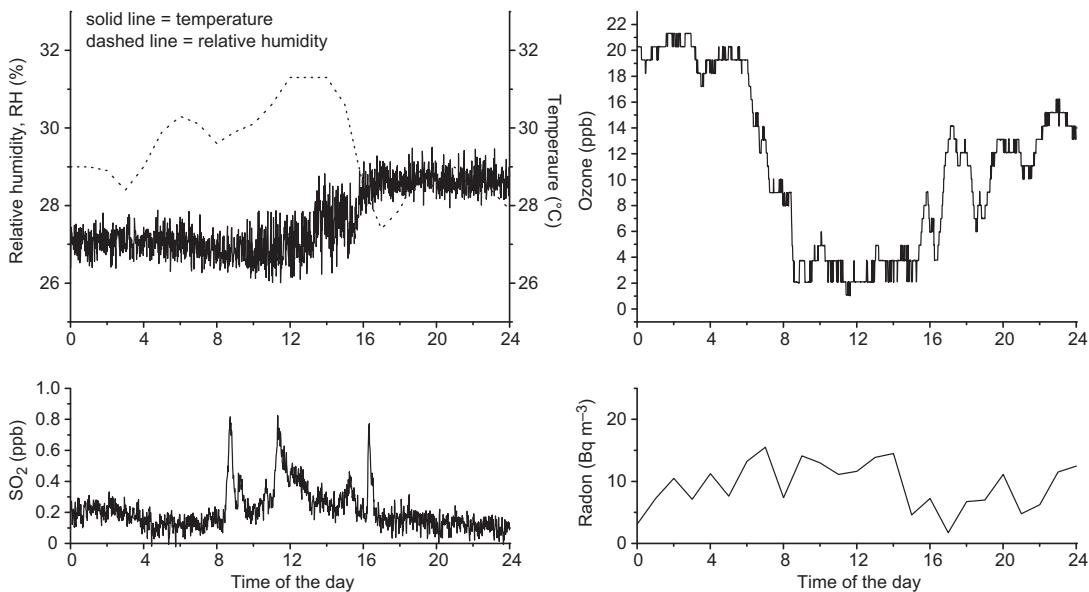


Fig. 2. Results of measurements on 7 April.

30 ppb even at night and the SO_2 concentration was around 0.2 ppb. These concentrations were similar to those in our experiment.

The lowest relative humidity found in the field during nocturnal events was 35%. The relative humidity in our experiment was slightly lower. Some recent studies on the role of humidity in nucleation events indicate that lower relative humidity in this range could favor the nucleation (Buzorius *et al.* 2001, Nilsson *et al.* 2001, Boy *et al.* 2002, Bonn *et al.* 2002, 2003, Hyvönen *et al.* 2005, Suni *et al.* 2008). So, the lower relative humidity in our laboratory experiment should lead to more, not less, frequent nucleation events than in the field.

The temperature during the experiment was 3 to 5 °C higher than the highest temperature found during the nocturnal events in Tumbarumba (23 °C). On the other hand, the concentration of eucalyptol in our experiments was 30 times higher than in the real atmosphere, so the

Table 1. Eucalyptol concentration during two hours after introducing it into the chamber.

Time	(ng l ⁻¹)	(cm ⁻³)
11:50	47.48	1.85×10^{11}
12:50	10.87	4.24×10^{10}
13:50	9.33	3.64×10^{10}

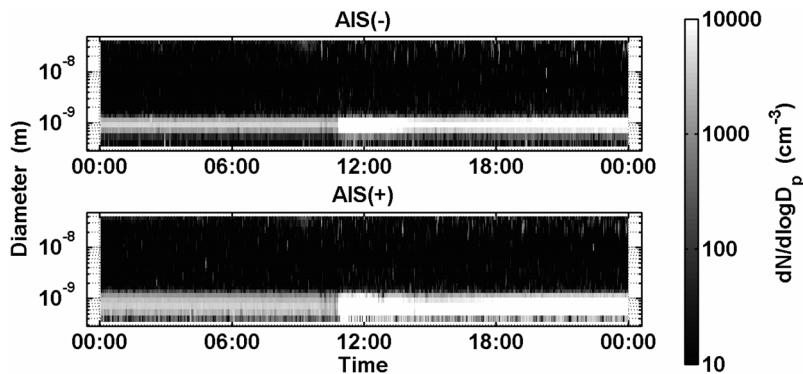


Fig. 3. Surface plot of air ion distribution measured on 8 April 2008 over size range of 0.34–40 nm in the Tammet diameter (Tammet *et al.* 1995)

negative effect caused by the higher temperature could have been compensated somewhat by the higher concentration of eucalyptol.

Radon experiment

To simulate the high concentration of radon observed in Tumbarumba (up to 100–200 Bq m⁻³ at night), we added radon in the chamber on 8 April (Fig. 3).

The air flow into the chamber via the tank containing the radon source was turned on and set up to 40 liters per minute at 10:46. One ml of eucalyptol was introduced into the chamber at the room temperature to obtain a higher concentration. At the start of the experiment, we observed a big burst of radon that had accumulated in the tank containing the radon source. The concentration reached a maximum of 2200 Bq m⁻³ and dropped then to around 500 Bq m⁻³ (Fig. 4). The eucalyptol concentrations was measured two times after the input of radon during five hours and the concentration was 20 and 25 ng l⁻¹ (7.8×10^{10} and 9.7×10^{10} cm⁻³).

The temperature increased to a maximum of 32 °C because the chamber door needed to be shut to keep radon inside the chamber, and the heat from the instruments warmed up the air quickly. As a result, also the relative humidity decreased to 25%. Ozone and SO₂ concentrations were around 8 ppb (1.97×10^{11} cm⁻³) and from 0.1 to 0.2 ppb (4.93 to 9.86×10^9 cm⁻³), respectively.

The presence of radon clearly increased the concentration of small ionic clusters. However,

it was not able to induce any nucleation events even with a high concentration of eucalyptol.

These two experiments showed that eucalyptol was not able to create any particles, even in the presence of radon. To understand better the reasons behind these results, we performed some quantum mechanical calculations to study the molecular mechanism behind these results in detail.

The oxidation rates of eucalyptol (table 1 in Calogirou *et al.* 1998) show that the absence of a double bond makes its oxidation by NO₃ and ozone quite difficult. The only oxidant that reacts with eucalyptol at a similar rate as with monoterpenes is the OH radical. However, since we are dealing with nocturnal events, this path has no relevance for our study. Consequently, we limited our calculations to the eucalyptol molecule.

Quantum chemical calculations

The most stable configuration for each cluster composition was obtained using the SIESTA program with the BLYP functional and the DZP basis set (Fig. 5). The structures have been drawn using the MOLDEN 3.8 visualization package (Schaftenaar *et al.* 2000). The values of ΔG were obtained (with respect to formation from free molecules) at 298 K and with monomer pressures of 1013 hPa for clusters containing only sulfuric acid and clusters containing one molecule of eucalyptol (Table 2).

If we compare the formation energies for clusters containing the same number of sulfuric acid molecules (n), we see how adding one molecule of eucalyptol leads to a more negative

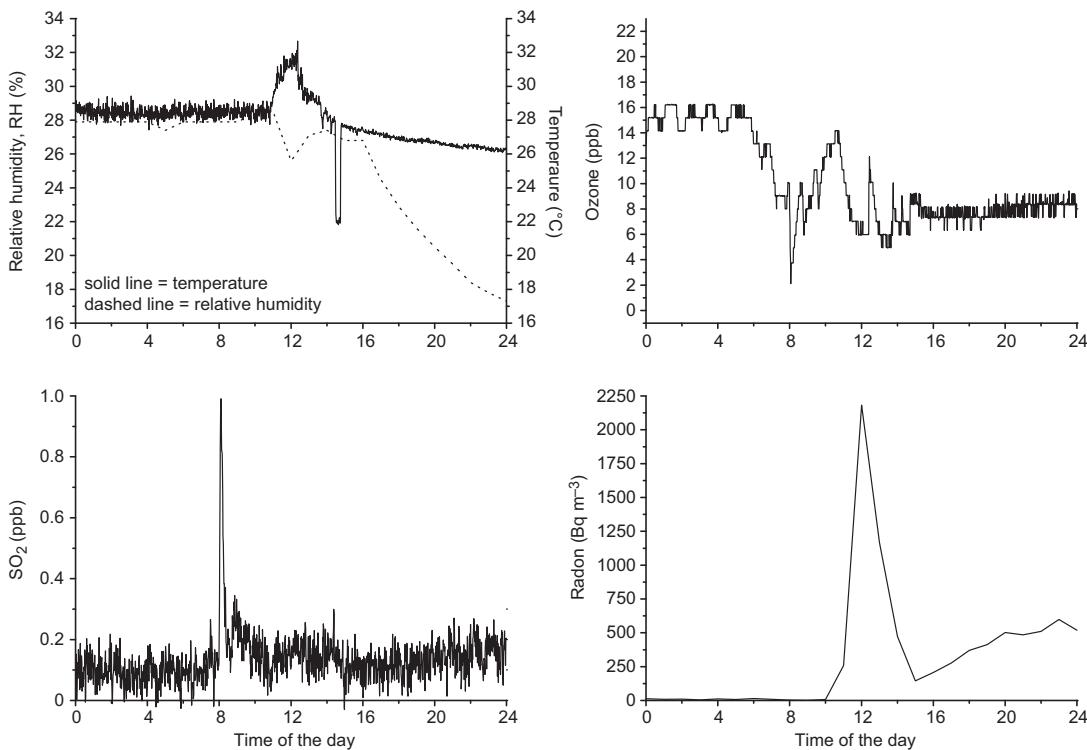


Fig. 4. Results of measurements on 8 April.

value of ΔG . The biggest difference was found for clusters containing three sulfuric acid molecules, for which the formation of clusters containing eucalyptol was around 7 kcal mol⁻¹ more favorable. For largest clusters containing four sulfuric acid molecules, this difference dropped to the value of 2 kcal mol⁻¹.

To understand the thermochemistry further, we also calculated the reaction free energies for the addition of one sulfuric acid molecule to the different clusters (Table 3). The addition of a second molecule of sulfuric acid to a mono-sulfuric acid cluster was slightly less favorable when a molecule of eucalyptol was present in the cluster (less than 1 kcal mol⁻¹). For clusters with two molecules of sulfuric acid, this effect was opposite: adding a third molecule of sulfuric acid was easier (around 2.5 kcal mol⁻¹) for clusters containing eucalyptol. For the addition of a fourth molecule for sulfuric acid, the energies for both clusters were virtually the same (difference of 0.21 kcal mol⁻¹).

The free energies for complexation at standard conditions (Table 2) cannot be directly used

to explain the nucleation-enhancing effect of eucalyptol. The free energy of cluster formation depends also on the relative atmospheric concentration of each molecular species in the atmosphere through the law of mass action (Ortega *et al.* 2008). So, the values of ΔG (Table 2) have to be corrected using atmospheric realistic concentrations. We chose a set of conditions to fit the conditions normally found in Tumbarumba

Table 2. Formation free energies at RI-CC2/Aug-cc-pV(T+d)Z level, temperature of 298 K and monomers pressure of 1013 hPa. The first column indicates the number of sulfuric acid molecules, the second column gives the free energy for a cluster consisting of sulfuric acid only, and the third column gives the free energy for a cluster with an additional eucalyptol molecule.

<i>n</i>	$(\text{H}_2\text{SO}_4)_n$ (kcal mol ⁻¹)	$\text{C}_{10}\text{H}_{18}\text{O} \times (\text{H}_2\text{SO}_4)_n$ (kcal mol ⁻¹)
1		-5.13
2	-6.18	-10.35
3	-9.58	-16.40
4	-12.78	-14.97

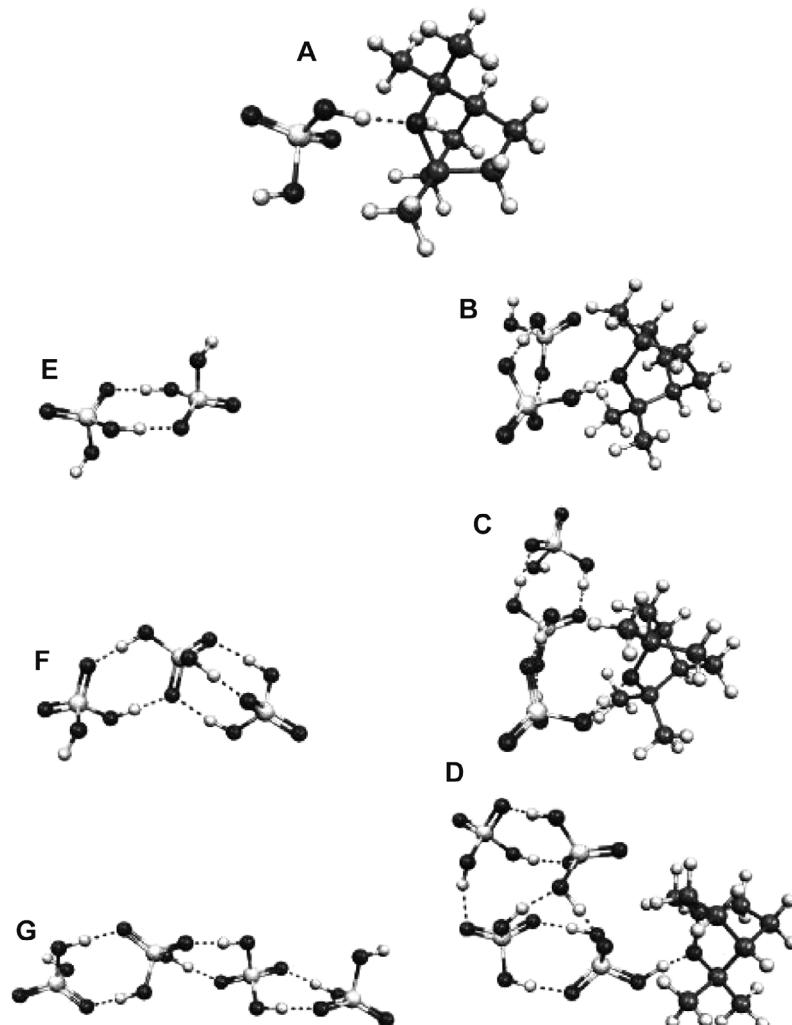


Fig. 5. Structures of the most stable cluster configurations obtained with the SIESTA model. **A**, **B**, **C** and **D** represent clusters with one molecule of eucalyptol and one to four molecules of sulfuric acid, respectively. **E**, **F** and **G** represent clusters containing from two to four sulfuric acid molecules, respectively. Black atoms correspond to oxygen, light grey to sulfur, dark grey to carbon and white to hydrogen.

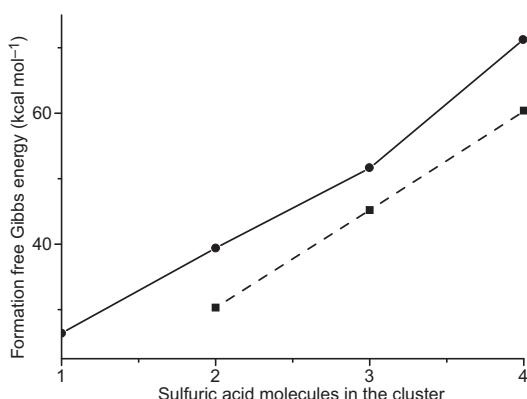


Fig. 6. Corrected values of ΔG versus number of sulfuric acid molecules in the cluster plots, dashed line corresponds to pure sulfuric acid clusters, and solid line corresponds to clusters containing one molecule of eucalyptol.

(Suni *et al.* 2008): 10^6 cm^{-3} of sulfuric acid calculated by a pseudo-steady state chemical box-model (Boy *et al.* 2005), 200 ppt ($4.9 \times 10^{10} \text{ cm}^{-3}$) of eucalyptol (corresponding roughly to the maximum concentration of eucalyptol found in Australia) and temperature of 298 K. These corrected values of ΔG (Table 4) can be presented as a function of the number of sulfuric acid molecules in the cluster (Fig. 6).

If we now take into account the corrected ΔG values obtained using the concentration of different species in the atmosphere (Table 4), the pure sulfuric acid clusters are more stable than clusters containing eucalyptol.

This difference is even more evident in the largest clusters we studied: the clusters contain-

Table 3. Formation free energy for addition of one sulfuric acid.

Reaction	ΔG (kcal mol ⁻¹)
$C_{10}H_{18}O + H_2SO_4 \rightarrow C_{10}H_{18}O \times H_2SO_4$	-5.13
$C_{10}H_{18}O \times H_2SO_4 + H_2SO_4 \rightarrow C_{10}H_{18}O \times (H_2SO_4)_2$	-5.22
$C_{10}H_{18}O \times (H_2SO_4)_2 + H_2SO_4 \rightarrow C_{10}H_{18}O \times (H_2SO_4)_3$	-6.04
$C_{10}H_{18}O \times (H_2SO_4)_3 + H_2SO_4 \rightarrow C_{10}H_{18}O \times (H_2SO_4)_4$	-3.41
$H_2SO_4 + H_2SO_4 \rightarrow (H_2SO_4)_2$	-6.18
$(H_2SO_4)_2 + H_2SO_4 \rightarrow (H_2SO_4)_3$	-3.40
$(H_2SO_4)_3 + H_2SO_4 \rightarrow (H_2SO_4)_4$	-3.20

ing four sulfuric acid molecules and one eucalyptol were less stable (11 kcal mol⁻¹) than those containing only sulfuric acid. The nucleation barrier was smaller for clusters containing only sulfuric acid than for clusters containing also eucalyptol (Fig. 6).

The results presented above are in a complete agreement with those obtained earlier in the experiments where we were not able to produce any events using eucalyptol. This observation is not surprising, considering that eucalyptol has only one oxygen atom available to interact with sulfuric acid molecules and that the position of the oxygen within the molecule makes this interaction even more difficult. A more oxidized molecule containing more functional groups, like a carboxylic acid, should be a better candidate for atmospheric nucleation than eucalyptol. Furthermore, the absence of a double bond in eucalyptol makes its oxidation by ozone impossible. Our conclusion is, therefore, that we have to look for the explanation of nocturnal events elsewhere, in other monoterpenes.

Conclusions

Eucalyptol is one of the most abundant organic compounds emitted from the eucalypt forest in Tumbarumba. The aim of this work was to study the role that eucalyptol could have in the nocturnal formation of new particles. We performed chamber experiments and quantum mechanic calculations to evaluate the importance of eucalyptol in this regard.

We were not able to induce any nucleation event using eucalyptol under conditions similar to those observed in Tumbarumba. Theoretical results indicated also that the eucalyptol cannot

Table 4. Formation free energies (kcal mol⁻¹) at the temperature of 298K, $[H_2SO_4] = 10^6$ cm⁻³ and $[C_{10}H_{18}O] = 200$ ppt. The first column indicates the number of sulfuric acid molecules, the second column gives the free energy for a cluster consisting of sulfuric acid only and the third column gives the free energy for a cluster with an additional eucalyptol molecule.

n	$(H_2SO_4)_n$	$C_{10}H_{18}O \times (H_2SO_4)_n$
1		26.4
2	30.3	39.4
3	45.2	51.6
4	60.3	71.3

stabilize sulfuric acid clusters. Hence, both experimental and theoretical studies led to the same conclusion: the presence of eucalyptol does not enhance the nucleation of sulfuric acid clusters. According to our results, the participation of eucalyptol in nocturnal events can be ruled out.

The explanation for the nocturnal events remains open. The most likely candidates for nocturnal events are VOCs containing one or more double bonds and having large enough oxidation rates. Therefore, further work in this direction is necessary: the work presented here should be repeated using different monoterpenes under dark conditions.

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References

- Ahlrichs R., Bär M., Häser J., Horn H. & Kölmel C. 1989. Electronic structure calculations on workstation comput-

- ers: the program system TURBOMOLE. *Chem. Phys. Lett.* 162: 165–169.
- Anttila T., Vehkamaki H., Napari I. & Kulmala M. 2005. Effect of ammonium bisulphate formation on atmospheric water-sulphuric acid-ammonia nucleation. *Boreal Env. Res.* 10: 511–523.
- Ball S.M., Hanson D.R., Eisele F.L. & McMurry P.H. 1999. Laboratory studies of particle nucleation: Initial results for H_2SO_4 , H_2O , and NH_3 vapors. *J. Geophys. Res.* 104: 23709–23718.
- Bonn B. & Moortgat K. 2003. Sesquiterpene ozonolysis: Origin of atmospheric new particle formation from biogenic hydrocarbons. *Geophys. Res. Lett.* 30: 1585–1588.
- Bonn B., Schuster G. & Moortgat G.K. 2002. Influence of water vapor on the process of new particle formation during monoterpene ozonolysis. *J. Phys. Chem. A* 106: 2869–2881.
- Bonn B., Kulmala M., Riipinen I., Sihto S.L. & Ruuskanen T.M. 2008. How biogenic terpenes govern the correlation between sulfuric acid concentrations and new particle formation. *J. Geophys. Res.* 113: D12209, doi:10.1029/2007JD009327.
- Boy M. & Kulmala M. 2002. Nucleation events in the continental boundary layer: influence of physical and meteorological parameters. *Atmos. Chem. Phys.* 2: 1–16.
- Boy M., Kazil J., Lovejoy E.R., Korhonen H., Guenther A. & Kulmala M. 2008. Relevance of ion-induced nucleation of sulfuric acid and water in the lower troposphere over the boreal forest at northern latitudes. *Atmos. Res.* 90: 151–158.
- Boy M., Kulmala M., Ruuskanen T.M., Pihlatie M., Reissell A., Aalto P.P., Kersten P., Dal Maso M., Hellen H., Hakola H., Jansson R., Hanke M. & Arnold F. 2005. Sulphuric acid closure and contribution to nucleation mode particle growth *Atmos. Chem. Phys.* 5: 863–878.
- Buzorius G., Rannik Ü., Nilsson E.D. & Kulmala M. 2001. Vertical fluxes and micrometeorology during aerosol particle formation events. *Tellus* 53B: 394–405.
- Calogirou A., Larsen B.R. & Kotzias D. 1998. Gas-phase terpene oxidation products: a review. *Atmos. Environ.* 33: 1423–1439.
- Dunning T.H.Jr., Peterson K.A. & Wilson A.K. 2001. Gaussian basis sets for use in correlated molecular calculations. X. The atoms aluminum through argon revisited. *J. Chem. Phys.* 114: 9244–9253.
- Hättig C. & Weigend F. 2000. CC2 excitation energy calculations on large molecules using the resolution of the identity approximation. *J. Chem. Phys.* 113: 5154–5161.
- Hoffmann T., Bandur R., Marggraf U. & Linscheid M. 1998. Molecular composition of organic aerosols formed in the α -pinene/ozone reaction: implications for new particle formation processes. *J. Geophys. Res.* 103: 25569–25578.
- Hyvönen S., Junninen H., Laakso L., Dal Maso M., Grönholm T., Bonn B., Kersten P., Aalto P., Hiltunen V., Pohja T., Launiainen S., Hari P., Mannila H. & Kulmala M. 2005. A look at aerosol formation using data mining techniques. *Atmos. Chem. Phys.* 5: 3345–3356.
- Junninen H., Hulkko M., Riipinen I., Nieminen T., Hirvikoski A., Suni T., Boy M., Lee S.-H., Vana M., Tammet H., Kerminen V.-M. & Kulmala M. 2008. Observation on night time growth of atmospheric clusters. *Tellus* 60B: 365–371.
- Korhonen P., Kulmala M., Laaksonen A., Viisanen Y., McGraw R. & Seinfeld J. 1999. Ternary nucleation of H_2SO_4 , NH_3 and H_2O in the atmosphere. *J. Geophys. Res.* 104: 26349–26353.
- Kulmala M., Pirjola L. & Mäkelä J.M. 2000. Stable sulphate clusters as a source of new atmospheric particles. *Nature* 404: 66–69.
- Kulmala M., Lehtinen K.E.J. & Laaksonen A. 2006. Cluster activation theory as an explanation of the linear dependence between formation rate of 3 nm particles and sulphuric acid concentration. *Atmos. Chem. Phys.* 6: 787–793.
- Kulmala M., Vehkamäki H., Petäjä T., Dal Maso M., Lauri A., Kerminen V.-M., Birmili W. & McMurry P.H. 2004a. Formation and growth rates of ultrafine atmospheric particles: a review of observations. *J. Aerosol. Sci.* 35: 143–176.
- Kulmala M., Suni T., Lehtinen K.E.J., Dal Maso M., Boy M., Reissell A., Rannik Ü., Aalto P., Kersten P., Hakola H., Bäck J., Hoffmann T., Vesala T. & Hari P. 2004b. A new feedback mechanism linking forest, aerosol, and climate. *Atmos. Chem. Phys.* 4: 557–562.
- Kulmala M., Riipinen I., Sipilä M., Manninen H.E., Petäjä T., Junninen H., Dal Maso M., Mordas G., Mirme A., Vana M., Hirsikko A., Laakso L., Harrison R.M., Hanson I., Leung C., Lehtinen K.E.J. & Kerminen V.-M. 2007. Toward direct measurement of atmospheric nucleation. *Science* 318: 89–92.
- Lovejoy E.R., Curtius J. & Froyd K.D. 2004. Atmospheric ion-induced nucleation of sulfuric acid and water. *J. Geophys. Res.* 109: D08204, doi:10.1029/2003JD004460.
- Miehlich B., Savin A., Stoll H. & Preuss H. 1989. Results obtained with the correlation-energy density functionals of Becke and Lee, Yang and Parr. *Chem. Phys. Lett.* 157: 200–206.
- Nilsson E.D., Paatero J. & Boy M. 2001. Effects of continental boundary layer evolution, convection, turbulence and entrainment, on aerosol formation. *Tellus* 53B: 462–478.
- O'Dowd C.D., Aalto P., Hämeri K., Kulmala M. & Hoffmann T. 2002. Atmospheric particles from organic vapours. *Nature* 416: 497–498.
- Ortega I.K., Kurtén T., Vehkamäki H. & Kulmala M. 2008. The role of ammonia in sulfuric acid ion induced nucleation. *Atmos. Chem. Phys.* 8: 2859–2867.
- Riipinen I., Sihto S.-L., Kulmala M., Arnold F., Dal Maso M., Birmili W., Saarnio K., Teiniälä K., Kerminen V.-M., Laaksonen A. & Lehtinen K.E.J. 2007. Connections between atmospheric sulphuric acid and new particle formation during QUEST III–IV campaigns in Hyytiälä and Heidelberg. *Atmos. Chem. Phys.* 7: 1899–1914.
- Schaftenaar G. & Noordik J.H. 2000. MOLDEN: a pre- and post-processing program for molecular and electronic structures. *J. Comput.-Aided Mol. Design.* 14: 123–134.
- Sihto S.-L., Kulmala M., Kerminen V.-M., Dal Maso M., Petäjä T., Riipinen I., Korhonen H., Arnold F., Janson R., Boy M., Laaksonen A. & Lehtinen K.E.J. 2006. Atmospheric sulfuric acid and aerosol formation: implications from atmospheric measurements for nucleation and early

- growth mechanisms. *Atmos. Chem. Phys.* 6: 4079–4091.
- Soler J.M., Artacho E., Gale J.D., Garcia A., Junquera J., Ordejon P. & Sanchez-Portal D. 2002. The SIESTA method for ab initio order-N materials simulation. *J. Phys.-Condens. Mat.* 14: 2745–2779.
- Suni T., Kulmala M., Hirsikko A., Bergman T., Laakso L., Aalto P.P., Luning R., Cleugh H., Zegelin S., Hughes D., Gorsel E., Kitchen M., Vana M., Hörrak U., Mirme A., Sevanto S., Twining K. & Tadros C. 2008. Formation and characteristics of ions and charged aerosol particles in a native Australian eucalypt forest. *Atmos. Chem. Phys.* 8: 129–139.
- Tammet H. 1995. Size and mobility of nanometer particles, clusters and ions. *J. Aerosol Sci.* 26: 459–475.
- Weber R.J., Marti J.J., McMurry P.H., Eisele F.L., Tanner D.J. & Jefferson A. 1996. Measured atmospheric new particle formation rates: Implications for nucleation mechanisms. *Chem. Eng. Commun.* 151: 53–64.
- Weber R.J., Marti J.J., McMurry P.H., Eisele F.L., Tanner D.J. & Jefferson A. 1997. Measurements of new particle formation and ultrafine particle growth rates at a clean continental site. *J. Geophys. Res.* 102: 4375–4385.
- Yu F. & Turco R. 2000. Ultrafine aerosol formation via ion-mediated nucleation, *Geophys. Res. Lett.* 27: 883–886.
- Yu F. & Turco R. 2008. Case studies of particle formation events observed in boreal forests: implications for nucleation mechanisms. *Atmos. Chem. Phys.* 8: 6085–6102.
- Yu F., Turco R.P. & Kärcher B. 1999. The possible role of organics in the formation and evolution of ultrafine aircraft particles. *J. Geophys. Res.* 104: 4079–4087.