

# Urban nitrous-oxide fluxes measured using the eddy-covariance technique in Helsinki, Finland

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Using the eddy covariance technique, nitrous oxide (N<sub>2</sub>O) fluxes were measured at the semi-urban SMEAR III station in Helsinki, Finland, between 21 June and 27 November 2012. The measurement period covered a seasonal change from summer to autumn allowing us to examine variations in the N<sub>2</sub>O fluxes by season. Also, varying land cover around the measurement site enabled us to study the effects of different urban land covers on the N<sub>2</sub>O exchange. Overall, the measurement surroundings acted as a source of N<sub>2</sub>O with a median and quartile deviation for the campaign of  $1.7 \pm 2.0 \mu\text{mol m}^{-2} \text{h}^{-1}$ . The net emissions were slightly higher from the direction of green areas than from the direction of roads under heavy traffic indicating that vegetation cannot be neglected in the urban N<sub>2</sub>O exchange studies. No seasonal change in the N<sub>2</sub>O flux during the measurement campaign was found.

## Introduction

Measured by its impact on radiative forcing, nitrous oxide (N<sub>2</sub>O) is the third important anthropogenic greenhouse gas in the atmosphere after carbon dioxide (CO<sub>2</sub>) and methane (IPCC 2013). Besides its warming effect on our climate (10% that of CO<sub>2</sub>), N<sub>2</sub>O also plays an important role in the chemical depletion of stratospheric ozone (Ravinshakara *et al.* 2009). Anthropogenic activities, including agriculture (16%), biomass burning (5%) and energy/industry (4%), account for approximately 40% of the global sources of N<sub>2</sub>O (Fowler *et al.* 2009). The rest of N<sub>2</sub>O is emitted mainly from oceans and soil. Due to anthropogenic activities, urban areas may act as hotspots

of N<sub>2</sub>O emissions not only from energy production and traffic, but also from managed vegetation. N<sub>2</sub>O emissions from vegetated surfaces are mainly related to the amount of nitrogen in soil and the microbial production of N<sub>2</sub>O via nitrification and denitrification processes. Thus, nitrogen-rich fertilizers play a key role in N<sub>2</sub>O emissions in managed green areas (Jones *et al.* 2011, Butterbach-Bahl *et al.* 2013). Soils can also act as a sink for N<sub>2</sub>O due to N<sub>2</sub>O consumption in denitrification or nitrifier denitrification processes under varying environmental conditions (Chapuis-Lardy *et al.* 2007).

In Helsinki, the three most important sources of N<sub>2</sub>O are energy production (43%), traffic (29%) and agriculture (13%) (Carney *et al.*

2009). These values, as well as greenhouse gas emission estimates in general, are calculated at an annual scale from national inventories. These inventories, however, lack information about the spatial and temporal variability of the emissions and sinks of the compounds inside a city, and typically ignore the effect of urban green areas. Variability of the exchange at a neighbourhood scale is important especially from the point of view of urban planning and climate mitigation.

The eddy covariance (EC) technique is a direct micrometeorological method to measure the exchange of greenhouse gases and other atmospheric pollutants between the surface and the atmosphere at a local scale, and it has successfully been used in various ecosystems for multiple compounds (e.g. Fowler *et al.* 2009, Aubinet *et al.* 2012). Only in the recent years, EC measurements of greenhouse gas emissions in urban areas have become popular, but the studies were mainly concentrated on CO<sub>2</sub> (e.g. Velasco and Roth 2010, Grimmond and Christen 2012). Studies have so far successfully compared these measurements with budget-based CO<sub>2</sub> net emission (e.g. Christen *et al.* 2011, Nordbo *et al.* 2012b), showing that the technique can successfully be used to follow the exchange of greenhouse gases in urban areas. To date, only one study has reported urban N<sub>2</sub>O fluxes from the city of Edinburgh, Scotland, over one month (Famulari *et al.* 2010). Thus, little is known about the spatial variability of N<sub>2</sub>O exchange at a city scale as well as its temporal variability both at the diurnal and seasonal scales. Urban areas are a mixture of built and green areas and therefore not only anthropogenic emissions, but also the effect of vegetation needs to be understood.

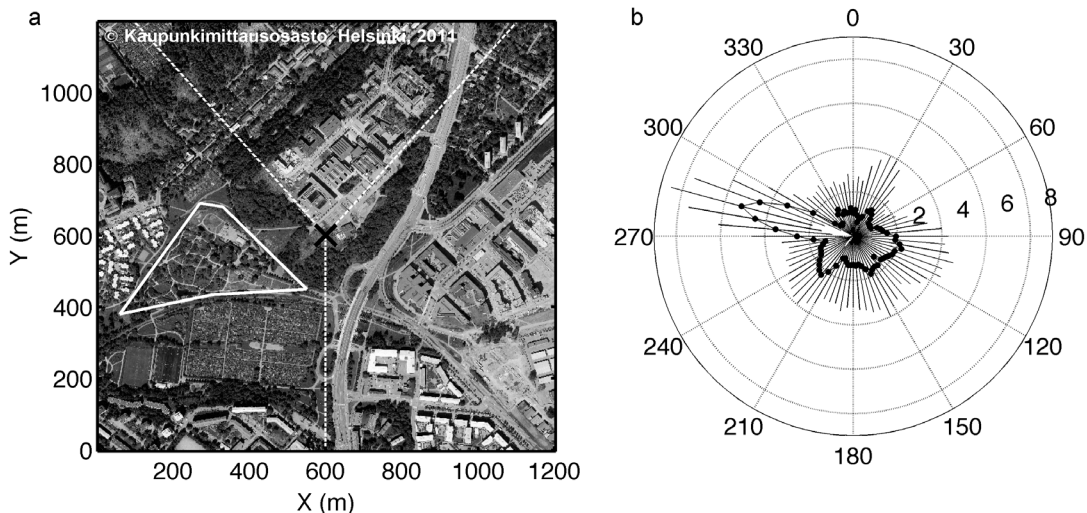
In this study, we report the results of an N<sub>2</sub>O flux EC measurement campaign carried out at the semi-urban SMEAR III (Station for Measuring Ecosystem–Atmosphere Relations, Järvi *et al.* 2009c) Kumpula site in Helsinki between June and November 2012. The measurement site is located at a high latitude and therefore is characterised by distinct seasons with varying vegetation activity. The surroundings of the flux-measurement tower fall into three surface-cover categories: roads under heavy traffic, buildings or green areas. This allowed us to examine the impact of different surface-cover types on the

exchange of N<sub>2</sub>O. Temporal changes in the N<sub>2</sub>O flux at the diurnal and seasonal scales were also examined, and correlations between net fluxes and environmental factors affecting the exchange were studied.

## Measurement site and measurements

### Measurement site

The measurements were carried out at the semi-urban measurement station SMEAR III Kumpula site (Fig. 1a) located 4 km NE from the Helsinki city centre (Järvi *et al.* 2009c). The eddy covariance measurements were performed from the top level of a 31-m-high lattice tower (60°12.17'N, 24°57.671'E) located on a hill, 26 m a.s.l. and 19–21 m above the surrounding terrain. The area around the tower can be divided into three surface cover sectors: built, road and vegetation, each representing the typical surface cover of the area (Vesala *et al.* 2008). The built sector in the northern direction (320°–40°) is dominated by university campus buildings and the Finnish Meteorological Institute (mean height 20 m) in the vicinity of the measurement tower. In this direction the observations can occasionally be in the roughness sublayer and therefore should be considered with caution. One of the main roads leading to Helsinki city centre passes through the road sector (40°–180°) with the closest distance between the road and the tower being 150 m. The area in-between is covered by deciduous forest with mainly birch, Norway maple, aspen, goat willow and bird cherry (Vesala *et al.* 2008). On the road, a typical workday traffic rate is around 44 000 vehicles per day (Lilleberg and Hellman 2011), and the vehicles have been found to be the main source of CO<sub>2</sub> and aerosol particle emissions in the area (Järvi *et al.* 2012, Ripamonti *et al.* 2013). In the vegetation sector (180°–320°), most of the surface is covered by green areas in the Kumpula Botanic Garden and in the City Allotment Garden. Fertilization and irrigation routines differ largely between the gardens. In the allotment garden, cultivation is only allowed between May and September, industrial fertilizers are banned and only manual irrigation is



**Fig. 1.** (a) Aerial photograph of the study site, where the flux tower is marked with black cross, the Kumpula Botanic Garden with white solid lines, and the different surface cover sectors with white dashed lines. (b) The median  $\text{N}_2\text{O}$  fluxes ( $\mu\text{mol m}^{-2} \text{h}^{-1}$ ) for  $20^\circ$  wind direction sectors calculated with  $5^\circ$  interval. Zero degree corresponds to north. The whiskers show the quartile deviations.

allowed. In the botanical garden, on the other hand, industrial fertilizers to keep the non-indigenous vegetation species alive are regularly used. In addition, when the botanic garden was rebuilt in 2006–2009, most of the plants and grass were planted on composted sewage sludge (L. Schulman, Finnish Museum of Natural History, University of Helsinki, pers. comm.), leaving lots of nitrogen (N)-rich organic material in the soil. Indigenous soils around the measurement tower are predominantly rock and till, and till or clay in low-lying areas, where both gardens are located.

### Measurement setup

The eddy covariance setup to measure the  $\text{N}_2\text{O}$  flux consisted of an ultrasonic anemometer (USA-1, Metek GmbH, Germany) and a TDL spectrometer (Tunable diode laser, TGA-100A, Campbell Scientific Inc., USA) to measure the  $\text{N}_2\text{O}$  mixing ratio. TDL comprises a single-mode diode laser controlled by current and temperature, and it is tuned to an infrared  $\text{N}_2\text{O}$  absorption band. The laser itself is installed in a liquid nitrogen dewar. A concentration measurement is obtained by passing the infrared laser beam through an absorption tube to the sample and

reference cells. Reference gas of 2200 ppm  $\text{N}_2\text{O}$  was drawn through the reference cell at the same temperature and pressure as the sample air in the sample cell. The TDL was calibrated twice during the measurement campaign and liquid nitrogen was added to the dewar bi-weekly. The conditions of the laser and the instrument were checked weekly, and the laser was adjusted twice during the campaign due to the line shift. The TDL spectrometer has successfully been used to measure the  $\text{N}_2\text{O}$  fluxes in various ecosystems (e.g. Wienhold *et al.* 1994, Pihlatie *et al.* 2005b, Rinne *et al.* 2005).

Complementary  $\text{CO}_2$  flux measurements were performed with a closed-path infrared gas analyzer (LI-7000, LI-COR, Lincoln, Nebraska, USA) using the same anemometer and partly the same inlet and tube. The anemometer was located 31 m above ground on a horizontal boom south-west from the tower. The air inlet of the gas analysers was situated 0.13 m below the anemometer centre. A 40-m-long steel tube with the inner diameter of 8 mm led to the  $\text{CO}_2$  analyser from where the air was further drawn to the TDL through a 10-m-long steel tube with the inner diameter of 4 mm. The flow rates in the tubes were  $20.2$  and  $17.2 \text{ l min}^{-1}$ , respectively, to ensure turbulent flow, and the sampling line

was heated with a power of  $16 \text{ W m}^{-2}$  to avoid condensation of water vapour. The frequency of the measurements was 10 Hz and the raw data were stored for post-processing. The  $\text{N}_2\text{O}$  flux measurements carried out between 21 June and 27 November 2012, whereas the  $\text{CO}_2$  flux is measured continuously.

## Auxiliary measurements

The auxiliary measurements cover global radiation (CNR1, Kipp&Zonen, Delft, Netherlands) and air temperature (Platinum resistant thermometer, Pt-100, built in-house) measured from the top boom of the EC tower at the height of 31 m. In addition, precipitation was measured on the roof of a nearby building using a weighting rain gauge (Pluvio2, Ott Messtechnik GmbH, Germany). Soil volumetric water content at 10 cm depth was measured continuously at two sites (three points each) in a neighbourhood called Viikki, around 4 km east from the Kumpula site. The measurements were made from growing media (area not covered by impervious surfaces) of street trees using Theta-probes (ML2x, Delta T Devices Ltd., Cambridge, UK) (Riikonen *et al.* 2011). Despite the distance between the EC flux tower and the soil-moisture measurement site, the results are considered to be representative for the changes in soil moisture also at the Kumpula site.

Hourly traffic rates were measured online 4 km from the Kumpula site by the City of Helsinki Planning Department. Observations were made hourly outside the rush hours (06:00–09:00 LST, Local Standard Time = UTC + 2 hours), and four times an hour at other times. These rates were converted to correspond to the traffic rates on the road next to the measurement site (Järvi *et al.* 2012).

## Data analysis and quality screening

### Flux calculations

The 30-min  $\text{N}_2\text{O}$  fluxes were calculated from the raw data using the maximum covariance technique with a fixed lag-time window (e.g. Foken

*et al.* 2012). Prior to the calculation of the final fluxes, raw data were de-spiked and linearly detrended, and a two-dimensional coordinate rotation was applied. The  $\text{N}_2\text{O}$  signal measured by the TDL was frequently dominated by a low-frequency signal drifting that previous studies have reported to be caused by the optical interference fringes. The drifting was suppressed by applying a high-pass autoregressive running mean filter with a time constant of 50 seconds before the flux calculations (e.g. Mammarella *et al.* 2010). As the sample air was dried prior to the analyser, and the measurement tube dampens the temperature fluctuations, no WPL correction was needed.

The  $\text{N}_2\text{O}$  fluxes ( $F_n$ ) were corrected for losses at the high- and low-frequency ends of the co-spectra as follows:

$$\frac{F_n^m}{F_n} = \frac{\int_0^{\infty} \text{TF}_H(f) \text{TF}_L(f) C_{ws}(f) df}{\int_0^{\infty} C_{ws}(f) df}, \quad (1)$$

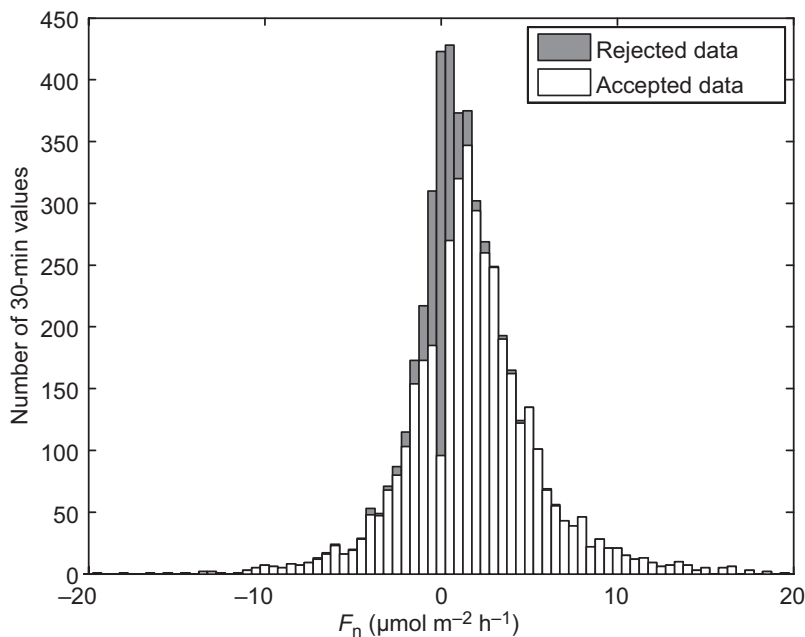
where  $F_n^m$  and  $F_n$  are the measured and un-attenuated  $\text{N}_2\text{O}$  fluxes ( $\mu\text{mol m}^{-2} \text{h}^{-1}$ ),  $f$  is the natural frequency (Hz),  $\text{TF}_H$  and  $\text{TF}_L$  are the transfer functions describing high- and low frequency losses. Both transfer functions were calculated as the product of appropriate theoretical transfer functions as described in Moncrieff *et al.* (1997) and Rannik and Vesala (1999).  $C_{ws}(f)$  is the site-specific temperature co-spectrum that is assumed to be un-attenuated. The stability dependency of  $C_{ws}(f)$  was taken into account when correcting each 30-minute  $F_n$  values. The peak co-spectral frequencies were taken as the maxima of fits to the co-spectra, and they were further converted to a normalized form as

$$n_m = f_m(z - z_d)/U, \quad (2)$$

where  $z$  is the measurement height,  $z_d$  is the displacement height (Vesala *et al.* 2008), and  $U$  is the wind speed. The normalized peak frequency is expected to be constant in unstable conditions and depend on atmospheric stability ( $\zeta$ ) in stable conditions with the form

$$n_m = \alpha_1 \left( 1 + \alpha_2 \zeta^{\alpha_3} \right), \quad (3)$$

where  $\alpha_1$ ,  $\alpha_2$ ,  $\alpha_3$  are fitting coefficients. Detailed



**Fig. 2.** Histogram of  $N_2O$  fluxes ( $F_n$ ) before and after the quality screening.

description of the data handling of the  $CO_2$  fluxes and spectral calculations is given in (Järvi *et al.* (2009a) and Nordbo *et al.* (2012a).

### Flux quality screening

During the 160-day measurement campaign, 36% of the  $N_2O$  flux data were lost due to power and instrument failures. In October, 98% of data were missing due to a data logging malfunction. Additionally, obvious outliers were removed from the  $N_2O$  flux data (covering 0.1%), and data were omitted according to the stationarity test that ensures that the stationarity criteria required by the EC technique is met (Foken and Wichura 1996). Data were omitted if the 30-minute flux values deviated more than 100% from the average of 5-minute subsections, corresponding to quality flag six (Foken *et al.* 2004). The stationarity test removed 10.7% of the  $N_2O$  fluxes (Fig. 2). The test mainly removes small fluxes increasing the campaign average by 16%. Finally, 53% of the  $N_2O$  flux data were available for further analysis.

Besides the data quality assurance, the flux detection limit and random error can be used to describe the quality of the flux measurements. The detection limit gives the minimal flux level that the instrument is able to detect due to uncer-

tainties arising from instrumental noise. It is defined as the standard deviation of the cross-correlation function far away from the peak, i.e., the maximum of the function, following Wienhold *et al.* (1994) with the modifications given in Nordbo *et al.* (2012a). The flux random error (Lenschow *et al.* 1994) gives the error due to a finite flux averaging time and it incorporates flux random error due to stochastic nature of turbulence as well as instrumental noise. Both the detection limit and the flux random error are converted to have units of covariance and normalized by the covariance values.

## Results

### Performance of the measurement setup

Quality of the measurements was assessed by inspecting the  $N_2O$  co-spectra. The median co-spectrum for unstable stratification follows a theoretical model (Kaimal *et al.* 1972) fairly well, with the exception of some deviation in the inertial sub-range (Fig. 3). In theory, white noise in the  $N_2O$  time series should not correlate with vertical wind speed and thus it should not be seen in a co-spectrum. Nevertheless, some deviation from the theoretical slope was found

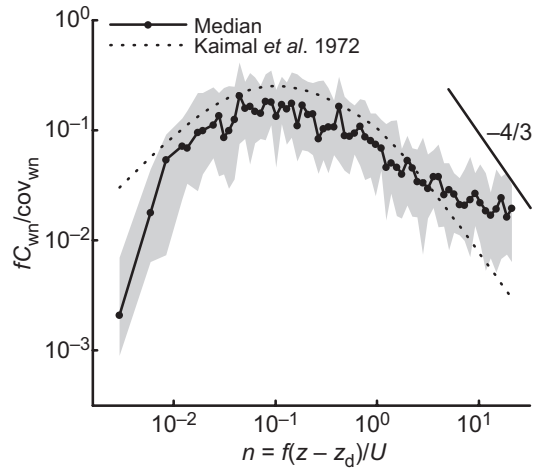
for a limited frequency interval at normalised frequencies above five. As this deviation is outside the energy-containing range, its impact on the flux will be small.

As expected, the co-spectral peak frequencies ( $n_m$ ) are constant in unstable situations with values  $0.10 \pm 0.04$  and  $0.08 \pm 0.04$  for  $\text{CO}_2$  and  $\text{N}_2\text{O}$ , respectively. On the other hand, a clear stability dependency of the  $\text{N}_2\text{O}$  peak frequency in the form  $n_m = 0.08(1 + 3.0\zeta^{0.47})$  was found (Eq. 3). The first fitting coefficient  $\alpha_1$  corresponds to the constant value in unstable conditions. Both the unstable mean values and the stability dependency in stable conditions are close to those found for the fluxes of sensible and latent heat,  $\text{CO}_2$  and aerosol particles at the same site (Vesala *et al.* 2008, Järvi *et al.* 2009b, 2009c).

The campaign median detection limit and the random error normalized with  $F_n$  were  $0.53 \pm 0.67$  and  $0.50 \pm 0.60$ , respectively, where the errors are quartile deviations (also hereinafter). Thus, on average the flux is twice the minimum flux that can be measured with the instrumentation, and it can be concluded that the instrumentation was sensitive enough for the measurements of low fluxes in a semi-urban environment. The relative detection limit and random uncertainty of the  $\text{N}_2\text{O}$  flux are over two times higher than obtained for momentum, sensible and latent heat, and  $\text{CO}_2$  fluxes at the same site during several years of measurements (Järvi *et al.* 2012, Nordbo *et al.* 2012a), and at an urban site in downtown Helsinki (Nordbo *et al.* 2013). The detection limit for  $\text{N}_2\text{O}$  fluxes from an agricultural soil has been determined to be 0.06–0.17 of the flux, which is 1/5 of that at our site (Wienhold *et al.* 1995).

## Weather conditions

During the measurement campaign, a seasonal change from summer to autumn took place (Fig. 4). The daytime (10:00–14:00) daily median global radiation ranged from  $\sim 830 \text{ W m}^{-2}$  to  $5 \text{ W m}^{-2}$  at the start and the end of the campaign, respectively, with the cloudy days clearly visible in the time series. Similarly, the daily air temperature was at its maximum  $23 \text{ }^\circ\text{C}$  in late July and the lowest daily temperature ( $-6 \text{ }^\circ\text{C}$ )



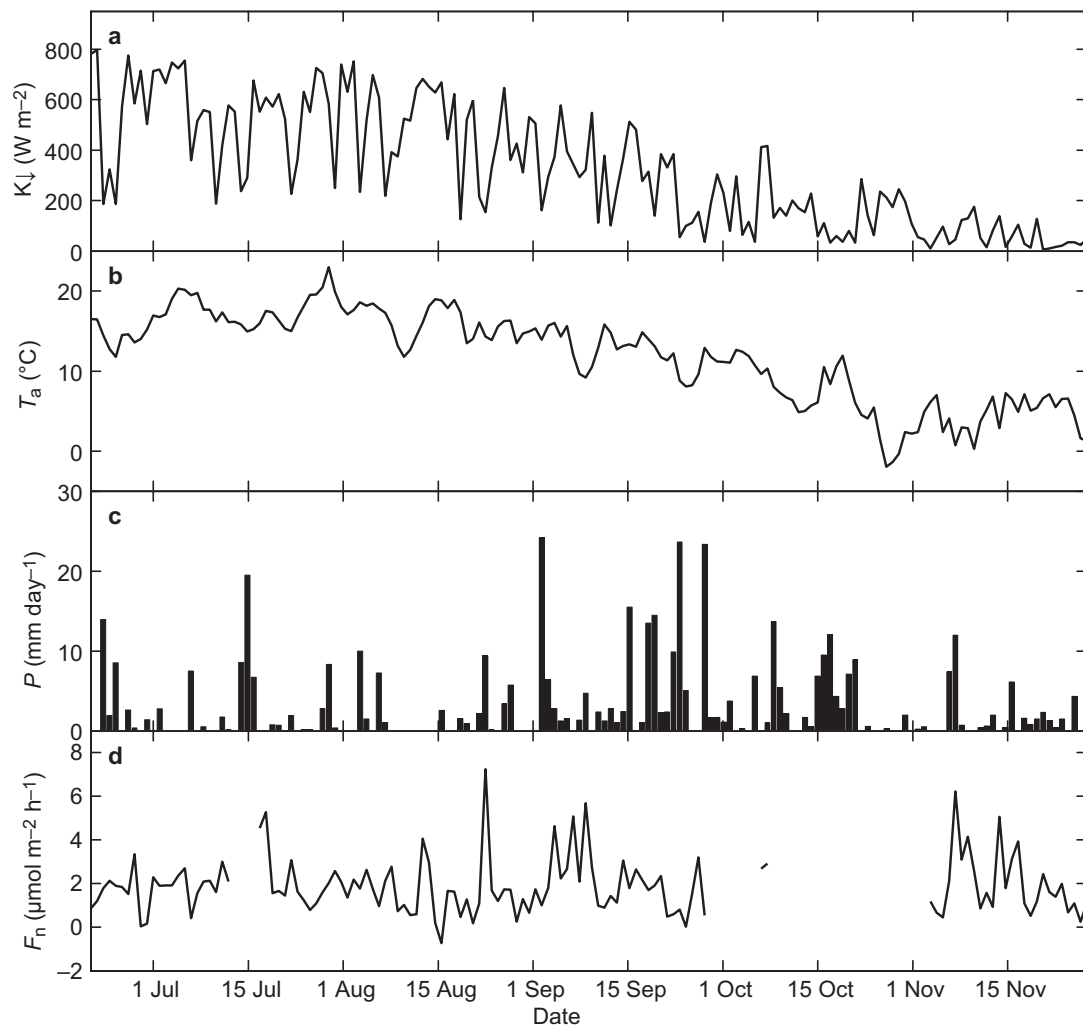
**Fig. 3.** Median  $\text{N}_2\text{O}$  flux co-spectrum ( $C_{wn}$ ) multiplied by the natural frequency ( $f$ ) and normalized with the covariance ( $\text{cov}_{wn}$ ) (black line), quartile deviation (gray area) and model co-spectrum by Kaimal *et al.* (1972) (dashed line) as a function of normalized frequency  $n$ . Data were binned into 76 logarithmically evenly-spaced classes and only co-spectra with  $F_n > 15 \mu\text{mol m}^{-2} \text{ h}^{-1}$  and unstable stratification were included (42 cases) in the plot.

was measured in early November. Precipitation was evenly distributed along the measurement campaign, but some separate weeks without rain were registered in 8–15 August and 22 Oct–2 Nov. During the measurement campaign, the most dominant wind direction was the S-SW, i.e., the direction where the highest fraction of vegetation cover was located (Fig. 5).

$F_n$  did not vary between summer and autumn. The maximum daily flux  $7.2 \mu\text{mol m}^{-2} \text{ h}^{-1}$  was found at the end of July. At the daily level, the surroundings acted as a source of  $\text{N}_2\text{O}$  with the exception in mid-August when slightly negative fluxes were measured (Fig. 4d).

## Spatial and seasonal variability of $\text{N}_2\text{O}$ fluxes

Median  $F_n$  ranged between 0.5 and  $4.6 \mu\text{mol m}^{-2} \text{ h}^{-1}$  with the highest upward fluxes from the direction  $270^\circ$ – $290^\circ$  (Fig. 1b). In this direction, the management area of the botanical garden as well as the newest part of the garden are located. These elevated fluxes can be explained by more intense fertilization, and thus more nitrogen rich soil, in the latter part of the



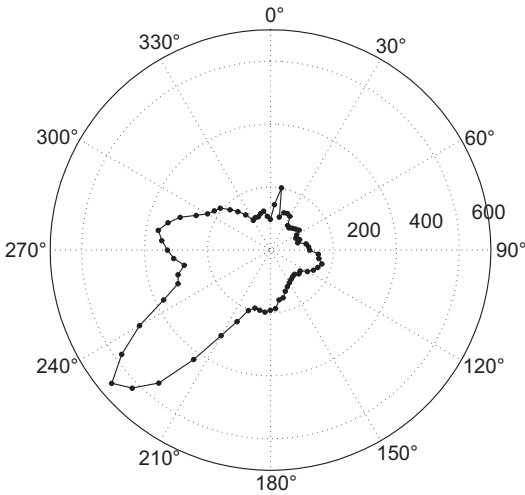
**Fig. 4.** Time series of daily mean (a) daytime (10:00–14:00) global radiation ( $K_{\downarrow}$ ), (b) air temperature ( $T_a$ ), (c) precipitation ( $P$ ), and (d)  $N_2O$  flux ( $F_n$ ) at the SMEAR III Kumpula site during 21 June–27 Nov 2012.

garden as compared with that in the older part of the garden and the nearby allotments. Also point  $N_2O$  emissions from the management areas, including fertilization storage and greenhouses, cannot be neglected. From the other parts of the garden and from the allotments ( $180^{\circ}$ – $260^{\circ}$ ), the fluxes were between 1 and  $2.5 \mu\text{mol m}^{-2} \text{h}^{-1}$ .

Another direction with elevated  $F_n$  was that of the road ( $80^{\circ}$ – $160^{\circ}$ ) where the fluxes reached  $2.2 \mu\text{mol m}^{-2} \text{h}^{-1}$ . Also the  $CO_2$  flux had a peak in that direction indicating the impact of traffic emissions on the  $N_2O$  signal (not shown). The lowest  $N_2O$  fluxes, on the other hand, were measured from the direction of the built sector

( $320^{\circ}$ – $20^{\circ}$ ) close to the measurement tower where both the amount of road traffic and vegetation are at their minima.

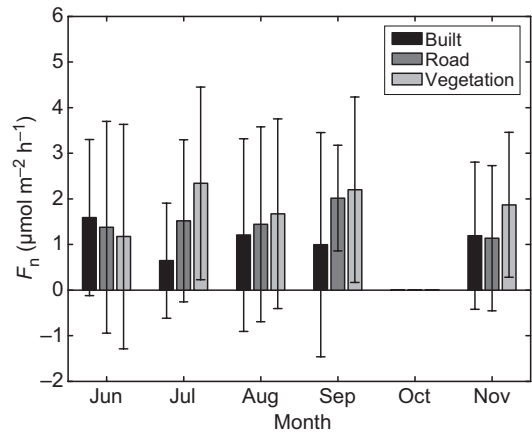
Based on the wind-direction dependence of  $F_n$ , it is reasonable to consider the different surface-cover sectors (built, road and vegetation) separately in further analysis. Excluding June, the highest  $N_2O$  emissions were recorded from the direction of the vegetation sector (Fig. 6). In June, only 9 days of data were available for analysis so the monthly median is not that representative. October was left empty due to the small amount of data caused by the instrument failure. The highest monthly emission,  $2.3 \mu\text{mol m}^{-2} \text{h}^{-1}$ ,



**Fig. 5.** The occurrence of 20° wind directions calculated for 5° interval; 0° corresponds to the north.

was measured in July from the vegetation sector, and the lowest  $0.6 \mu\text{mol m}^{-2} \text{h}^{-1}$  in same month from the built sector. When averaged over the whole measurement campaign, the median fluxes for vegetation, road and built areas were 2.0, 1.4 and  $1.2 \mu\text{mol m}^{-2} \text{h}^{-1}$ , respectively (Table 1).

During the measurement campaign,  $F_n$  did not change seasonally (Fig. 4) despite the fact that there is more road traffic outside summer months in addition to the increasing need for heating in autumn. Generally in Helsinki, and also around the flux tower, most buildings use district heating, but some houses north of the flux tower may additionally use fireplaces. In summer,  $F_n$  is more affected by vegetation and soil whereas in autumn litter may have stimu-



**Fig. 6.** Median  $\text{N}_2\text{O}$  fluxes ( $F_n$ ) for the difference surface cover sectors (built, road and vegetation) in each month. Quartile deviations are shown with error bars.

lated soil  $\text{N}_2\text{O}$  emissions (Pihlatie *et al.* 2007). Median  $F_n$  during the measurement campaign was  $1.7 \pm 2.0 \mu\text{mol m}^{-2} \text{h}^{-1}$ . The median net flux in Edinburgh in November was tenfold ( $11.7 \mu\text{mol m}^{-2} \text{h}^{-1}$ , Famulari *et al.* 2010) than measured at our site in the same month. This can be explained by the proximity of a dense downtown area to their flux tower, where there is likely more traffic than on the roads passing nearby our measurement site. Unfortunately, no other  $F_n$  values from urban areas using the same micrometeorological techniques have been reported and the impact of dense urban areas on the fluxes cannot be verified. Our median value is slightly smaller than obtained for agricultural sites, which can be explained by intensive nitrogen fertilization of agricultural soils leading to

**Table 1.** Monthly median daytime (10:00–14:00) global radiation ( $K\downarrow$ ) and air temperature ( $T$ ), monthly cumulative precipitation ( $P$ ) and median  $\text{N}_2\text{O}$  fluxes ( $F_n$ ) for each surface cover sector around the EC flux tower. Quartile deviations of  $F_n$  are shown for the campaign medians.

	Jun	Jul	Aug	Sep	Oct	Nov	Campaign
$K\downarrow$ ( $\text{W m}^{-2}$ )	723	609	531	218	121	33	254
$T$ ( $^{\circ}\text{C}$ )	13.3	17.5	16.1	12.7	6.8	4.6	12.7
$P$ ( $\text{mm month}^{-1}$ )	90	63	47	175	100	65	540
$F_n$ ( $\mu\text{mol m}^{-2} \text{h}^{-1}$ )							
Built	1.59	0.64	1.20	0.99	–	1.19	$1.18 \pm 1.94$
Road	1.38	1.52	1.44	2.01	–	1.14	$1.41 \pm 1.89$
Veg	1.17	2.34	1.67	2.20	–	1.87	$1.96 \pm 1.87$
All	1.30	1.93	1.53	2.11	–	1.63	$1.70 \pm 1.97$



higher  $N_2O$  emissions (Table 2). As compared with that for more natural ecosystems, our value is slightly higher. On the other hand, around 100 times higher emissions were measured from a landfill in Helsinki using the same instrument ( $193 \mu\text{mol m}^{-2} \text{h}^{-1}$ ; *see Rinne et al. 2005*).

### Diurnal behaviour of the $N_2O$ fluxes

As the flux tower was rarely downwind from the built sector (Fig. 5), lots of data for this direction are missing.

Although,  $F_n$  showed highly variable behaviour similarly to what was found by Famulari *et al.* (2010), but still a distinct diurnal cycles were seen both in the vegetation and road sectors (Fig. 7a). Downwind from these two directions, however, the behaviour was different indicating differences in sources/sinks. Using the Mann-Whitney  $U$ -test we found that the differences between the fluxes in these two surface-cover sectors were significant ( $U = 1980$ ,  $p = 0.0109$ ). In the vegetation sector, maximum  $F_n$  ( $3.9 \mu\text{mol m}^{-2} \text{h}^{-1}$ ) was measured in the afternoon, whereas the nocturnal fluxes remained below  $2 \mu\text{mol m}^{-2} \text{h}^{-1}$ . This indicates that the  $N_2O$  emissions were related to soil microbial activity, which in general increases with increasing soil temperature. Also, as the maximum  $F_n$  occurred at the same time with a maximum uptake of  $CO_2$  ( $-7 \mu\text{mol m}^{-2} \text{s}^{-1}$ ), emissions may also be related to the vegetation activity as  $N_2O$  can be emitted from trees via transpiration (Pihlatie *et al.* 2005a).

Slightly higher maximum  $F_n$  than that in the vegetation sector was measured for the road sector during daytime ( $4.2 \mu\text{mol m}^{-2} \text{h}^{-1}$ ).  $F_n$  started to increase around 05:00 LST corresponding well with the start of the morning rush hour (Fig. 6c). At the same time, an increase was also seen in the  $CO_2$  flux,  $F_c$  (Fig. 7b). The morning and rush-hour maxima of  $F_n$  were  $3.5 \mu\text{mol m}^{-2} \text{h}^{-1}$  and  $4.1 \mu\text{mol m}^{-2} \text{h}^{-1}$ , respectively, and they also coincided with the  $F_c$  maxima ( $9.7$  and  $10.4 \mu\text{mol m}^{-2} \text{s}^{-1}$ , respectively). Thus,  $F_n$  and  $F_c$  showed traffic-related similarity in the direction of the road under heavy traffic.

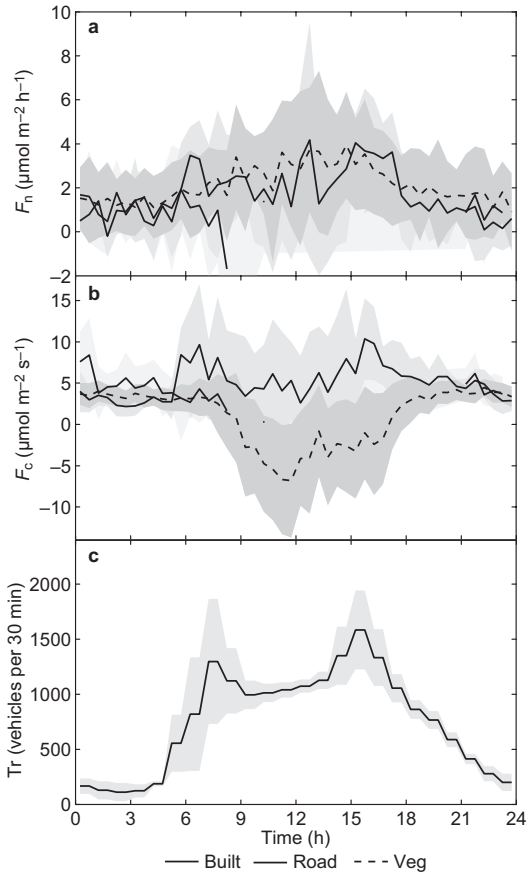
A linear regression between the median fluxes and traffic values showed a clear dependence of  $F_n$  on  $F_c$  (RMSE =  $0.94 \mu\text{mol m}^{-2} \text{h}^{-1}$ ,  $r = 0.59$ ) and traffic rate (RMSE =  $0.76 \mu\text{mol m}^{-2} \text{h}^{-1}$ ,  $r = 0.76$ ) (Fig. 8). The intercept of the latter regression equation ( $0.27 \mu\text{mol m}^{-2} \text{h}^{-1}$ ; *see Fig. 8b*) can be considered to represent contributions from other sources than traffic to the  $N_2O$  signal, which in this direction is mainly vegetation in residential areas. This would correspond well with  $F_n$  measured in natural/agricultural ecosystems in Europe (Table 2).

Our diurnal peak fluxes were much lower than  $32 \mu\text{mol m}^{-2} \text{h}^{-1}$  reported for Edinburgh in Famulari *et al.* (2010). In order to understand the differences, we drew a plot similar to their fig. 6b, where median  $F_n$  for the road sector were calculated for 200 vehicles per half-hour bins (Fig. 9). Fluxes increased with increasing vehicle numbers, but the maximum fluxes were found for the range 1200–1400 vehicles per 30 min not for the maximum traffic rates.

**Table 2.** Mean  $N_2O$  fluxes ( $F_n$ ,  $\mu\text{mol m}^{-2} \text{h}^{-1}$ ) measured using the eddy-covariance technique in different environments. Reference values were chosen according to the representativeness of ecosystems and environments.

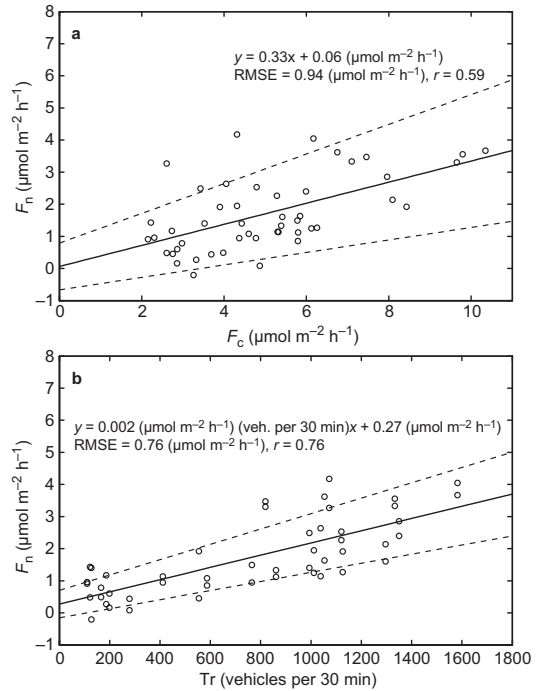
Location	Site	Period	$F_n$	Reference
Helsinki, Finland	Landfill	Aug 2003	192.7*	Rinne <i>et al.</i> (2005)
Edinburgh, UK	Urban	Nov–Dec 2005	11.7**	Famulari <i>et al.</i> (2010)
Helsinki, Finland	Urban	Jun–Nov 2012	1.7**	This study
Harford, NY, USA	Agricultural	Jun–Jul 2008	2.8*	Molodovskaya <i>et al.</i> (2011)
Lägeren, Switzerland	Mixed forest	Oct–Nov 2005	0.8*	Eugster <i>et al.</i> (2007)
Easter Bush, UK	Grassland	Jan–Dec 2002	0.5**	Fletcher <i>et al.</i> (2007)
		Jan–Dec 2003	0.2**	
Lille Bøgeskov, Denmark	Beech forest floor	May–Jun 2005	0.4**	Pihlatie <i>et al.</i> (2005b)
Kalevansuo, Finland	Peatland forest	Apr–Jun 2007	0.3**	Mammarella <i>et al.</i> (2010)

\* mean, \*\* median.



**Fig. 7.** Diurnal behaviour of the median (a)  $N_2O$  flux ( $F_n$ ), (b)  $CO_2$  flux ( $F_c$ ), and (c) traffic rates ( $Tr$ ) during the measurement campaign. Fluxes are shown for the built, road and vegetation sectors. The shaded areas show the quartile deviations. Only data when both  $F_n$  and  $F_c$  were available were considered and only half hours with more than 10 data points are shown in the plot. This way, only nocturnal data from the built surface cover sectors are plotted.

Similar behaviour was also found for particle fluxes at the site (Järvi *et al.* 2009b), and it is thought to be related to the amount of congestion, increased number of heavy-duty vehicles during the rush hours, and the enhanced mixing of the atmosphere as these traffic rates correspond to the daytime hours. The overall dependency in Helsinki and Edinburgh is qualitatively similar with increasing fluxes with higher traffic, but the flux magnitudes are systematically one order of magnitude smaller in Helsinki than in Edinburgh. Since in our case, the road next to the flux tower is known to be the main pollutant

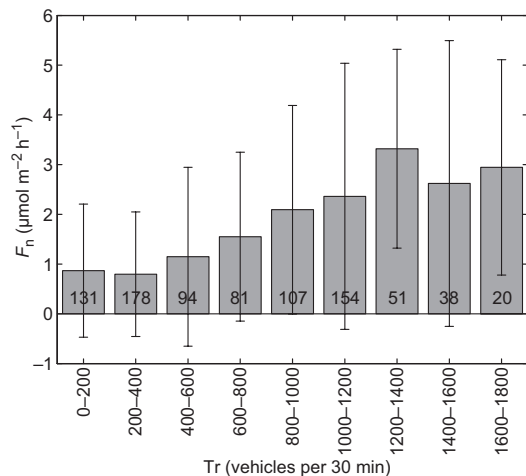


**Fig. 8.** Correlations between the median half-hour  $N_2O$  fluxes ( $F_n$ ) and (a)  $CO_2$  flux ( $F_c$ ), and (b) traffic rate ( $Tr$ ). The data correspond to the 48 data points representing the road sector in Fig. 6. The solid line is the linear regression and the dashed lines indicate the 95% confidence limits.

emitter in that direction, we assume that the different relationship in Edinburgh is caused by the downtown area where several roads with more traffic are located. Similarly to  $F_n$ , also  $F_c$  was found to be greater (three times in November) in Edinburgh than at our site in Helsinki (Nemitz *et al.* 2002).

### $N_2O$ emissions factors relative to $CO_2$ emissions

It is common to analyse the emission ratio (ER) of different air pollutants respective to  $F_c$ , as  $CO_2$  is considered to be a good tracer for anthropogenic emissions (e.g. Järvi *et al.* 2009b, Famulari *et al.* 2010, Vogt *et al.* 2011). ER was thus calculated as a campaign median for the road sector yielding  $6.4 \times 10^{-5}$ . Obtained ER corresponds well with the annual value of  $6.7 \times 10^{-5}$  obtained from emission inventories for residential and



**Fig. 9.** Dependence of half-hourly  $\text{N}_2\text{O}$  fluxes ( $F_n$ ) on local traffic rate ( $\text{Tr}$ ) at the SMEAR III Kumpula road sector. Median values for bins of 200 vehicles per 30 min were calculated. Error bars show the quartile deviations and the numbers in the bars indicate how many half-hour data points were in each bin.

transport emissions in Helsinki (Carney *et al.* 2009). The emissions from the direction of the road can be considered to be mainly a combination of these two sources, but unfortunately it was not possible to determine their relative contributions, which causes discrepancy in the comparison of ER from inventories and EC measurements.

ER obtained from the measurements can be used as a proxy for annual  $\text{N}_2\text{O}$  emissions with the aid of annual  $\text{CO}_2$  emissions from the road sector. Järvi *et al.* (2012) reported a five-year mean annual  $\text{CO}_2$  emission of  $3500 \text{ g C m}^{-2}$  for the road sector, which would give  $0.26 \text{ g N m}^{-2}$  for annual  $\text{N}_2\text{O}$  emission. This value is only approximate as unfortunately we do not have information on how  $\text{N}_2\text{O}$  fluxes behave in winter and spring at our measurement site.

### Impact of soil moisture on $\text{N}_2\text{O}$ emissions

$\text{N}_2\text{O}$  emissions from soils are known to be dependent on soil moisture and its changes as water affects the oxygen availability for soil microbes (e.g. Pihlatie *et al.* 2005b, Butterbach-Bahl *et al.* 2013). Therefore, we analysed our

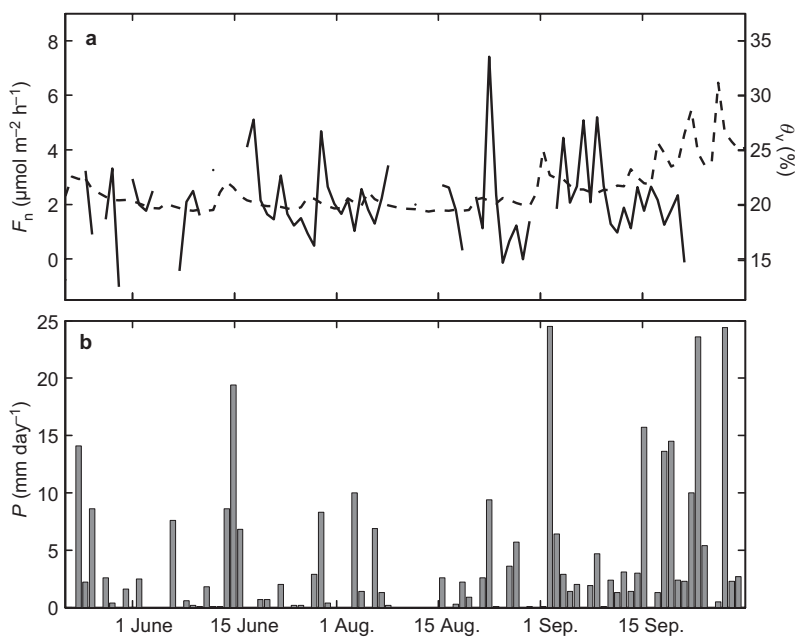
$\text{N}_2\text{O}$  fluxes from the direction of the vegetation sector together with the average volumetric soil water content ( $\theta_v$ , %; see Fig. 10). Wetting of the soil followed precipitation events well (e.g. on 15–17 July and 2 September) despite the 4-km distance between our site and the site where precipitation was measured. This proves that these precipitation values can be used as a reference for soil moisture changes in Kumpula.

Most of the time, the increase and decrease in daily  $F_n$  followed the changes in soil moisture well (Fig. 10). Almost always an increase in soil moisture was followed by a peak in  $\text{N}_2\text{O}$  emissions. However, the correlation between the daily means did not exist ( $r = 0.08$ ), but this could be expected as the vegetation around the flux tower is more intensively managed than at the soil-moisture measurement sites. However, the comparison suggests that also at urban sites the  $\text{N}_2\text{O}$  net fluxes are, to some degree, dependent on the soil moisture and not only on anthropogenic emissions. This implies that microbial activity in the soil and vegetation play an important role in  $\text{N}_2\text{O}$  emissions at urban sites.  $F_n$  were also compared with other meteorological variables, including air temperature and global radiation, but no dependency was found.

### Conclusions

In Helsinki, an  $\text{N}_2\text{O}$  flux measurement campaign using the eddy-covariance technique was carried out in June–November 2012. The measurement period covered the seasonal change from summer to autumn allowing us to examine the seasonal variation in  $\text{N}_2\text{O}$  fluxes. Furthermore, the different surface-cover sectors around our flux tower allowed us to examine the effects of different urban land-covers on the measured fluxes.

Overall, the measurement surrounding was a source of  $\text{N}_2\text{O}$  with a median net flux of  $1.7 \pm 2.0 \mu\text{mol m}^{-2} \text{h}^{-1}$ . Higher fluxes were measured from the highly vegetated area (median =  $2.0 \mu\text{mol m}^{-2} \text{h}^{-1}$ ) than from the direction of road under heavy traffic (median =  $1.4 \mu\text{mol m}^{-2} \text{h}^{-1}$ ). The highest emissions were recorded from the new area of the Kumpula Botanic Garden, where the vegetation is more managed as compared



**Fig. 10.** Daily mean (a)  $\text{N}_2\text{O}$  flux ( $F_n$ ) at the SMEAR III vegetation sector (solid line) and soil volumetric water content (dashed line) measured in Viikki ( $\theta_v$ ), and (b) precipitation ( $P$ ) between 21 June and 31 September 2012.

to the other green areas around the flux tower. In the area of high fraction of vegetation,  $\text{N}_2\text{O}$  emissions seemed to be dependent on the soil moisture similarly to non-urban ecosystems. In the direction of the road, a clear effect of local traffic on the  $\text{N}_2\text{O}$  fluxes in a diurnal cycle was found and the correlation between the  $\text{N}_2\text{O}$  and  $\text{CO}_2$  fluxes, and between the  $\text{N}_2\text{O}$  flux and traffic rate were high (0.59 and 0.76, respectively; see Fig. 8). Using the measurements, we also calculated emission ratios for the road sector ( $6.4 \times 10^{-5}$ ) that matched well with the ratios obtained from an inventory study. The  $\text{N}_2\text{O}$  fluxes did not show a seasonal behaviour during the autumn transformation.

The results suggest that the eddy-covariance technique can be used to measure the exchange of  $\text{N}_2\text{O}$  at our site, and that the impact of urban vegetation on the  $\text{N}_2\text{O}$  exchange. So far only two urban  $\text{N}_2\text{O}$  studies, including the current one, have been made. The three-fold difference in the values of  $F_n$  emphasizes the need for additional observations in urban areas that differ in their population density and degree of development. These are needed in order to draw general conclusions about  $\text{N}_2\text{O}$  exchange in urban areas and dependencies of the emissions on road traffic and vegetation.

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